

## Analysis for Skin SST using the GEOS model and GSI: preliminary results<sup>1</sup>

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Oct. 10, 2012

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<sup>1</sup>Acknowledgement: We thank John Derber & Xu Li for the NSST code in GSI

# Outline

1. Motivation
2. Introduction to vertical temperature profile & modeling
3. GSI Analysis for skin temperature
4. Experiment setting
5. Results
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# Motivation

Accurate modeling and estimation of air-sea fluxes requires accurate representation of SST (Saunders, 1967; Fairall et al., 1996).

Ocean surface vertical Temp. is highly variable within a day (Donlon et al., 2002, Ward, 2006)

## Current Status...

- ▶ AGCMs typically use a weekly averaged blend of in situ measurements (depth of few cm – meters) and satellite retrievals (Reynolds et al., 2002)
- ▶ OGCMs consider SST to be the 'mean' temperature of the surface of the ocean (for e.g., top layer of the model, typically 1–5 m thick)

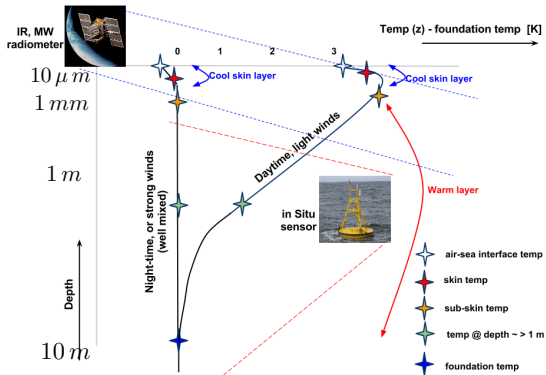
## Our Objectives:

1. Model the diurnal variation of SST (Takaya et al., 2010 and Zeng & Beljaars, 2005)
2. Assimilate for the vertical profile of temperature:  $T(z)$
3. Estimate the sea surface skin temperature ( $T_s$ ); Use  $T_s$  to compute air-sea fluxes

Ultimately

improved **surface modules**  $\rightarrow$  **AGCM**  $\oplus$  **OGCM**  $\Rightarrow$  **coupled assimilation**

# Introduction- $T(z)$ profile



See GHR SST for details:

<https://www.ghrsst.org/ghrsst-science/sst-definitions>

# Introduction- Modeling $T(z)$ profile

- ▶ **Cool-skin layer**: In the upper surface (few mm) molecular (thermal) diffusion is dominant. It is due to combined effects of **net longwave radiation, sensible and latent heat fluxes**
- ▶ **Warm layer**: mm to few meters thick, follows diurnal warming.
- ▶ Cool layer is always persistent (day & night)
- ▶ Warm layer builds up due to solar heating
- ▶ Clouds, strong winds (erode stratification)
- ▶ Warm layer can persist even into night or next day (Gentemann et al., 2009)
- ▶ Turbulent diffusion dominates molecular diffusion effects in the warm layer

IR & MW sat. retrievals → temperature decrease due to cool-layer

in Situ measurements → diurnal warming w.r.t. foundation temperature

Our  $T(z)$  profile is based on

- ▶ diurnal warming model of Zeng & Beljaars, 2005
- ▶ cool-skin layer of Fairall et al., 1996

# GSI Analysis for skin temperature

- Read in additional background fields to compute temp.  $T(z_{ob})$  at any  $z_{ob}$
- Observation (or penetration) depth ( $z_{ob}$ ) for radiance measurements

**IR** AVHRR, AIRS, HIRS, IASI, Sounders, Imagers, etc,  $\rightarrow 15\mu m$

**MW** AMSUA, AMSUB, MHS, SSMI, AMSRE, etc,  $\rightarrow 1.25mm$

in Situ observations have  $z_{ob}$  recorded

- Given  $z_{ob}$ , time and location (lat, lon) of the observation, background temperature is computed as

$$T(z_{ob}) = \begin{cases} T_d + dt_w - (1 - \frac{z_{ob}}{\delta})dt_c & \text{if } 0 \leq z_{ob} < \delta \rightarrow \text{COOL LAYER} \\ T_d + dt_w - [\frac{z_{ob} - \delta}{z_w - \delta}]^\mu dt_w & \text{if } \delta \leq z_{ob} < z_w \rightarrow \text{WARM LAYER} \end{cases} \quad (1)$$

- CRTM surface requires specification of the Jacobian  $\frac{\partial T(z)}{\partial T_s}$ , it is currently set to 1

$\Rightarrow$  Entire  $T(z)$  profile shifts to accomodate the change in  $T_s$

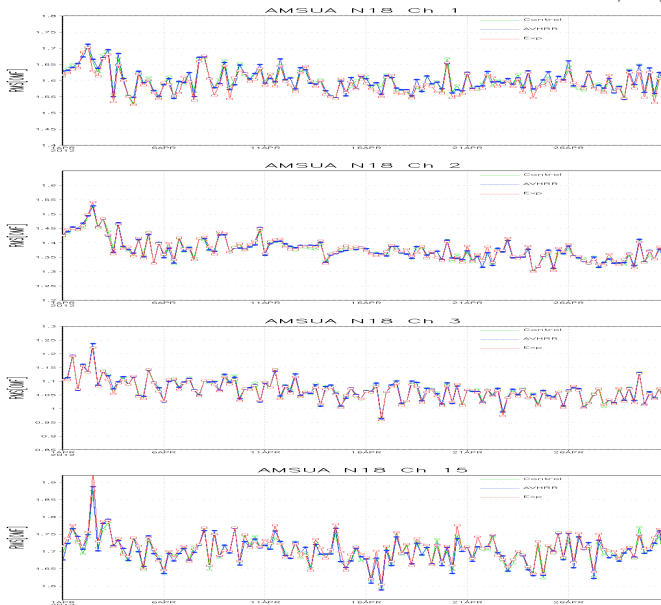
- Analysis writes out analyzed  $T_s^{ana}$  and the model applies  $T_s^{ana} - T_s^{bkg}$  increment through an IAU mechanism

# Experiment Setting

Exp. Name	AVHRR	$T_s$ Analysis
Control	<input type="checkbox"/>	<input type="checkbox"/>
AVHRR	<input checked="" type="checkbox"/>	<input type="checkbox"/>
Exp	<input checked="" type="checkbox"/>	<input checked="" type="checkbox"/>

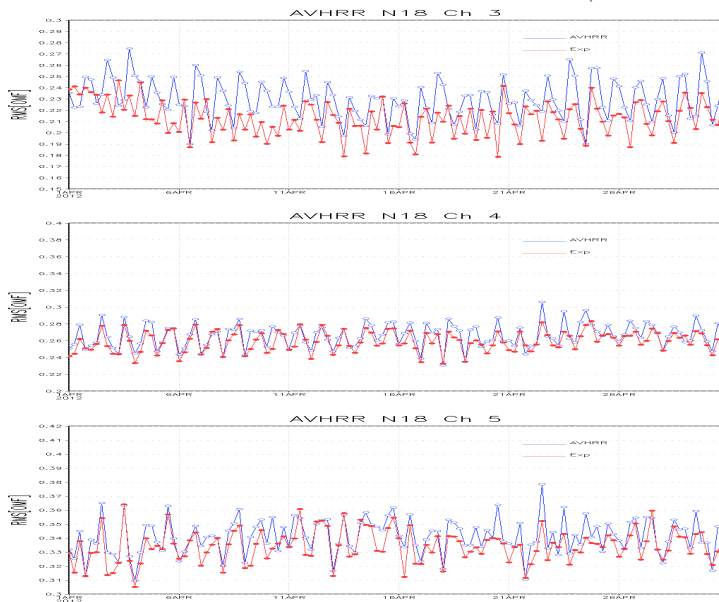
- ▶ Resolution:  $2^\circ \times 2.5^\circ$
- ▶ Foundation SST and sea-ice concentration are from weekly Reynolds
- ▶ All experiments have an active diurnal model for skin temperature computation
- ▶ DAS cycling experiment for April 2012
- ▶ *Additional* relevant observations are from AVHRR

# Results- RMS of OMF AMSUA N18- Ch: 1, 2, 3 & 15

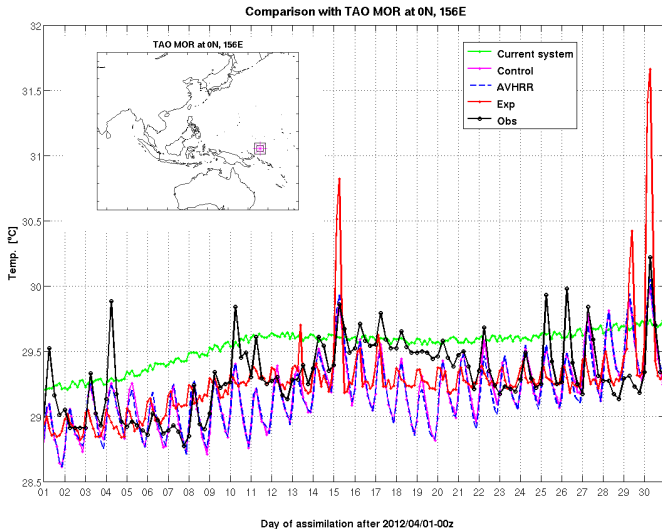




# Results- RMS of OMF AVHRR N18- Ch 3, 4 & 5

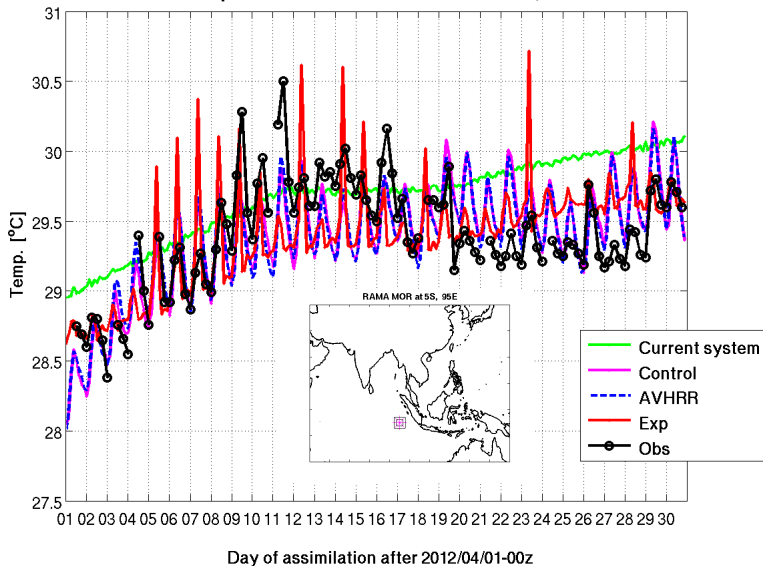


# Results- Comparison with in Situ buoy measurements



# Results- Comparison with in Situ buoy measurements

## Comparison with RAMA MOR at 5S, 95E



# Summary & work in progress

1. Accurate modeling & estimation of skin temperature requires modeling and analysis of  $T(z)$  profile
2. Within the GEOS we now model the vertical temperature profile including diurnal warming & cool-skin layer
3.  $T_s$  analysis components using the GEOS & GSI are ready
4. Work is in progress to use in Situ observations (*modsbufr*) in analysis
5. Currently we are evaluating and tuning performance for various observing systems

# References

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